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J. Phys. A: Math. Theor. **41** (2008) 145206 (17pp)

doi:10.1088/1751-8113/41/14/145206

# Solitary wave and periodic wave solutions for the thermally forced gravity waves in atmosphere

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Received 17 November 2007, in final form 17 February 2008 Published 26 March 2008 Online at stacks.iop.org/JPhysA/41/145206

# Abstract

By introducing a new transformation, a new direct and unified algebraic method for constructing multiple travelling wave solutions of general nonlinear evolution equations is presented and implemented in a computer algebraic system, which extends Fan's direct algebraic method to the case when r > 4. The solutions of a first-order nonlinear ordinary differential equation with a higher degree nonlinear term and Fan's direct algebraic method of obtaining exact solutions to nonlinear partial differential equations are applied to the combined KdV-mKdV-GKdV equation, which is derived from a simple incompressible non-hydrostatic Boussinesq equation with the influence of thermal forcing and is applied to investigate internal gravity waves in the atmosphere. As a result, by taking advantage of the new first-order nonlinear ordinary differential equation with a fifth-degree nonlinear term and an eighthdegree nonlinear term, periodic wave solutions associated with the Jacobin elliptic function and the bell and kink profile solitary wave solutions are obtained under the effect of thermal forcing. Most importantly, the mechanism of propagation and generation of the periodic waves and the solitary waves is analysed in detail according to the values of the heating parameter, which show that the effect of heating in atmosphere helps to excite westerly or easterly propagating periodic internal gravity waves and internal solitary waves in atmosphere, which are affected by the local excitation structures in atmosphere. In addition, as an illustrative sample, the properties of the solitary wave solution and Jacobin periodic solution are shown by some figures under the consideration of heating interaction.

PACS numbers: 02.30.Jr, 04.20.Jb

(Some figures in this article are in colour only in the electronic version)

1751-8113/08/145206+17\$30.00 © 2008 IOP Publishing Ltd Printed in the UK

### 1. Introduction

Gravity waves in the atmosphere are a subject of broad interest and play a significant role in weather and climate, such as the rainstorm of typhoons, orographic precipitation and atmospheric circulation [1, 2]. One important mechanism of gravity wave production and propagation in the atmosphere is that the airflow response to a transient heat source, which plays a key role in the generation and propagation of internal gravity waves in the atmosphere [3, 4]. One fundamental object of research on wave phenomena is to search for travelling wave solutions [5–11]. Many methods to construct exact solutions of nonlinear wave equations have been established and developed, such as the Bäcklund transformation [12–14], Hirota's bilinear method [15–17], tanh-function method [18–24], extended tanh-function method [25, 26], variational iteration methods [27, 28], collocation method [29–31], Adomian Padé approximation [32], inverse scattering method [13, 33], Darboux transformation [34–38] and so on. Very recently, the F-expansion [39, 40], auxiliary equation [41–43], Fan's sub-equation [44, 45], modified extended Fan's sub-equation methods [46], which are straightforward and effective, were proposed for constructing periodic wave solutions for some nonlinear evolution equations.

The aim of this paper is to consider the first-order nonlinear ordinary differential equation (ODE) with a higher degree nonlinear term

$$\frac{\mathrm{d}\varphi}{\mathrm{d}\xi} = \epsilon \sqrt{\sum_{j=0}^{r} c_j \varphi^j} \tag{1.1}$$

and with a lower degree nonlinear term

$$\frac{\mathrm{d}\phi}{\mathrm{d}\xi} = \epsilon \sqrt{\sum_{j=0}^{s} c_j \phi^j},\tag{1.2}$$

where  $\epsilon = \pm 1, r > 4, s < r$  and  $c_j$  (j = 0, 1, ..., r) are constants. In fact, when we make a transformation

when we make a transformation
$$\varphi \to \phi^{(s-2)/(r-2)} \tag{1.3}$$

for equation (1.1) and next specify the coefficients  $c_j$  (j = 0, 1, ..., r), then equation (1.1) is reduced to equation (1.2).

For example, based on Fan's sub-equation method a first-order ODE with a fourth-degree nonlinear term [47–54] is considered, namely

$$\frac{d\varphi}{d\xi} = \epsilon \sqrt{c_0 + c_1 \varphi + c_2 \varphi^2 + c_3 \varphi^3 + c_4 \varphi^4},$$
(1.4)

where we choose s = 4 in equation (1.3). If we make a transformation

$$\varphi \to \phi^{2/(r-2)},$$
 (1.5)

then equation (1.1) is reduced to equation (1.4). If r = 5, when we make a transformation  $\varphi \rightarrow \phi^{2/3}$ , then the following first-order ODE with a fifth-degree nonlinear term

$$\frac{d\varphi}{d\xi} = \epsilon \sqrt{c_0 + c_1 \varphi + c_2 \varphi^2 + c_3 \varphi^3 + c_4 \varphi^4 + c_5 \varphi^5}$$
(1.6)

can be reduced to equation (1.4).

The rest of this paper is organized as follows: in section 2, we derive the mathematical model from the incompressible non-hydrostatic Boussinesq equation, which governs the internal gravity waves in the atmosphere; in section 3, we apply the auxiliary equation method to find various periodic and solitary wave solutions of internal waves, which is used to explain the effect of heat on gravity waves in the atmosphere; in section 4, some conclusions are given.

### 2. Derivation of the KdV-mKdV-GmKdV equation

Consider the following two-dimensional nonlinear incompressible non-hydrostatic Boussinesq equation [55, 56], which consists of the nonlinear horizontal momentum equation, mass continuity equation and the nonlinear thermodynamic equation:

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + \varpi \frac{\partial u}{\partial z} = -\frac{1}{\rho} \frac{\partial p'}{\partial x}$$
(2.1*a*)

$$\frac{\partial \varpi}{\partial t} + u \frac{\partial \varpi}{\partial x} + \varpi \frac{\partial \varpi}{\partial z} = -\frac{1}{\rho} \frac{\partial p'}{\partial z} + B$$
(2.1*b*)

$$\frac{\partial \rho'}{\partial t} + u \frac{\partial \rho'}{\partial x} - \rho \varpi \frac{N^2}{g} = Q$$
(2.1c)

$$\frac{\partial u}{\partial x} + \frac{\partial \varpi}{\partial z} = 0, \tag{2.1d}$$

where u and  $\overline{\omega}$  are the velocities, N is the buoyancy frequency per mass unit, p' represents the perturbation pressure and  $B = -g \frac{\rho'}{\rho}$  denotes the buoyancy with  $\rho$  denoting the density, gdenoting the gravitational acceleration and  $\rho'$  representing the perturbation density.

Introducing the transformations

$$u(x, z, t) = u(\xi), \qquad \varpi(x, z, t) = \varpi(\xi), \qquad p'(x, z, t) = p(\xi), \frac{\rho'(x, z, t)}{\rho(x, z, t)} = \pi(\xi), \qquad \xi = kx + nz - wt$$
(2.2)

where w represents an angular frequency, and  $\vec{K} = (k, n)$  denotes the wave vector, we can rewrite equation (2.1) as

$$(-w + ku + n\varpi)\frac{\mathrm{d}u}{\mathrm{d}\xi} = -\frac{k}{\rho}\frac{\mathrm{d}p}{\mathrm{d}\xi}$$
(2.3*a*)

$$(-w + ku + n\varpi)\frac{\mathrm{d}\varpi}{\mathrm{d}\xi} = -\frac{n}{\rho}\frac{\mathrm{d}p}{\mathrm{d}\xi} - g\pi$$
(2.3b)

$$(-w+ku)\frac{\mathrm{d}\pi}{\mathrm{d}\xi} - \frac{N^2}{g}\varpi = Q \tag{2.3c}$$

$$k\frac{\mathrm{d}u}{\mathrm{d}\xi} + n\frac{\mathrm{d}\varpi}{\mathrm{d}\xi} = 0. \tag{2.3d}$$

Integrating equation (2.3*d*) with respect to  $\xi$  once and taking the integration constant as zero, one can get

$$\varpi = -\frac{ku}{n}.\tag{2.4}$$

Substitution of equation (2.4) into equation (2.3) yields

$$\frac{d^2u}{d\xi^2} + \frac{k^2 N^2 u}{w(k^2 + n^2)(w - ku)} - \frac{kngQ}{\rho w(k^2 + n^2)(w - ku)} = 0.$$
 (2.5)

If  $u \ll \frac{w}{k}$ , then  $F(u) = \frac{1}{w-ku}$  can be expanded as

$$F(u) = \frac{1}{w - ku} = \frac{1}{w} \left( 1 + \frac{ku}{w} + \frac{k^2 u^2}{w^2} + \frac{k^3 u^3}{w^3} + \cdots \right).$$
 (2.6)

Neglecting high power terms in the polynomial of F(u), if we choose that

$$\frac{d^2u}{d\xi^2} + \frac{k^2 N^2 u}{w^2 (k^2 + n^2)} \left( 1 + \frac{ku}{w} + \frac{k^2 u^2}{w^2} + \frac{k^3 u^3}{w^3} \right) - \frac{kng Q}{\rho w^2 (k^2 + n^2)} \left( 1 + \frac{ku}{w} + \frac{k^2 u^2}{w^2} \right) = 0, \quad (2.7)$$
  
then we can reduce equation (1.5) to  
$$\frac{d^2u}{\omega} + \frac{k^5 N^2 u^4}{w^2} + \frac{k^4 N^2 u^3}{\omega} + \frac{k^3 (N^2 \rho \omega - Qgn) u^2}{w^2}$$

$$\frac{d\xi^{2}}{d\xi^{2}} + \frac{\omega^{5}(k^{2} + n^{2})}{\omega^{5}(k^{2} + n^{2})} + \frac{\omega^{4}(k^{2} + n^{2})}{\omega^{4}(k^{2} + n^{2})\rho} + \frac{k^{2}(N^{2}\rho\omega - Qgn)u}{\omega^{3}(k^{2} + n^{2})\rho} - \frac{Qgnk}{\omega^{2}(k^{2} + n^{2})\rho} = 0.$$
(2.8)

Differentiating equation (2.8) once with respect to  $\xi$  leads to

$$\frac{d^{3}u}{d\xi^{3}} + \frac{4k^{5}N^{2}u^{3}}{\omega^{5}(k^{2}+n^{2})}\frac{du}{d\xi} + \frac{3k^{4}N^{2}u^{2}}{\omega^{4}(k^{2}+n^{2})}\frac{du}{d\xi} + \frac{2k^{3}(N^{2}\rho\omega - Qgn)u}{(k^{2}+n^{2})\rho\omega^{4}}\frac{du}{d\xi} + \frac{k^{2}(N^{2}\rho\omega - Qgn)}{\omega^{3}(k^{2}+n^{2})\rho}\frac{du}{d\xi} = 0,$$
(2.9)

which is the ordinary differential equation that the generalized combined KdV-mKdV-GmKdV equation corresponds to.

# 3. Solutions to the KdV-mKdV-GmKdV equation

Let us now focus our attention on equation (2.9) and introducing the transformations, equation (2.9) is generated as

$$\frac{\mathrm{d}^3 u}{\mathrm{d}\xi^3} + (c + \alpha u + \beta u^2 + \gamma u^3) \frac{\mathrm{d}u}{\mathrm{d}\xi} = 0, \qquad (3.1)$$

in which

$$\alpha = \frac{2k^{3}(\rho\omega N^{2} - ngQ)}{(k^{2} + n^{2})\rho\omega^{4}}, \qquad \beta = \frac{3k^{4}N^{2}}{\omega^{4}(k^{2} + n^{2})},$$
  

$$\gamma = \frac{4k^{5}N^{2}}{\omega^{5}(k^{2} + n^{2})}, \qquad c = \frac{k^{2}(\rho\omega N^{2} - ngQ)}{\rho\omega^{3}(k^{2} + n^{2})}.$$
(3.2)

If we let

$$u = \sum_{i=1}^{n} a_i \varphi^i \tag{3.3}$$

with  $\varphi$  satisfying

$$\frac{\mathrm{d}\varphi}{\mathrm{d}\xi} = \epsilon \sqrt{\sum_{j=0}^{r} c_j \varphi^j},\tag{3.4}$$

we can conclude that r = 3n + 2 through balancing terms  $u^3 \frac{du}{d\xi}$  and  $\frac{d^3u}{d\xi^3}$  in equation (3.1). Not losing generality, when we choose n = 1 and r = 5, we can set  $u = a_0 + a_1\varphi$  with  $\varphi$  satisfying

$$\frac{d\varphi}{d\xi} = \epsilon \sqrt{c_0 + c_1 \varphi + c_2 \varphi^2 + c_3 \varphi^3 + c_4 \varphi^4 + c_5 \varphi^5}$$
(3.5)

and when we take n = 2 and r = 8, we can set  $u = a_0 + a_1\varphi + a_1\varphi^2$  with  $\varphi$  satisfying

$$\frac{\mathrm{d}\varphi}{\mathrm{d}\xi} = \epsilon \sqrt{c_0 + c_1 \varphi + c_2 \varphi^2 + c_3 \varphi^3 + c_4 \varphi^4 + c_5 \varphi^5 + c_6 \varphi^6 + c_7 \varphi^7 + c_8 \varphi^8}.$$
 (3.6)

### 3.1. The first-order ODE with a fifth-degree nonlinear term

With the aid of Maple9, substitution of equation (3.5) into equation (3.1) shows that the set of algebraic equation possesses the solution

$$c_{5} = -\frac{\gamma a_{1}^{3}}{10}, \qquad c_{4} = -\frac{\gamma a_{0} a_{1}^{2}}{2} - \frac{\beta a_{1}^{2}}{6}, \qquad c_{2} = -\alpha a_{0} - c - \gamma a_{0}^{3} - \beta a_{0}^{2},$$

$$c_{3} = -\frac{2\beta a_{0} a_{1}}{3} - \frac{\alpha a_{1}}{3} - \gamma a_{0}^{2} a_{1},$$
(3.7)

where  $a_1 \neq 0$ ,  $a_0$ , c and  $c_0$  are arbitrary constants.

Considering the seven parameters  $a_i$  (i = 0, 1) and  $c_j$  (j = 0, 1, ..., 5), especially the  $\gamma$  parameter, we present some types of solutions in the following cases.

*Case 1.* If  $\gamma = 0$ , then equation (3.1) is reduced to the well-known combined KdV–mKdV equation, and equation (3.5) takes in the form

$$\frac{d\varphi}{d\xi} = \epsilon \sqrt{c_0 + c_1 \varphi + c_2 \varphi^2 + c_3 \varphi^3 + c_4 \varphi^4}.$$
(3.8)

If we choose  $a_0 = -\frac{\alpha}{2\beta}$  in equation (3.7), then  $c_4 \neq 0$ ,  $c_3 = 0$ ; we have the following solutions to equation (3.1):

$$u_1 = -\frac{\alpha}{2\beta} + \frac{\sqrt{6(\alpha^2 - 4\beta c)}}{2\beta} \operatorname{sech}\left(\sqrt{\frac{\alpha^2 - 4\beta c}{4\beta}}\xi\right),\tag{3.9}$$

where it requires that  $\alpha^2 - 4\beta c > 0$  and  $\beta > 0$ , so these solutions must satisfy the condition

$$n^{2}g^{2}Q^{2} - 2\rho^{2}w^{2}N^{4} + \rho w N^{2}ngQ > 0, \qquad N^{2} > 0, \qquad (3.10)$$

which tells us that in the stable atmosphere  $(N^2 > 0)$ , heating (Q > 0) helps to generate upward-propagating waves, but cooling (Q < 0) helps to generate downward-propagating waves.

From equation (2.4), we have that the slope of lines of constant phase in the x-z plane is

$$\frac{\mathrm{d}z}{\mathrm{d}x} = -\frac{k}{n}.\tag{3.11}$$

Considering equations (3.10) and (3.11), it is not difficult to see that in the stable atmosphere ( $N^2 > 0$ ), heating (Q > 0) helps to generate westerly propagating waves if the slope of lines of constant phase is positive, and heating (Q > 0) helps to generate easterly propagating waves if the slope of lines of constant phase is negative.

However, in the stable atmosphere  $(N^2 > 0)$ , cooling (Q < 0) helps to generate easterly propagating waves if the slope of lines of constant phase is positive and cooling (Q < 0) helps to generate westerly propagating waves if the slope of lines of constant phase is negative. Figure 1 shows the propagating behaviour in qualitative terms

$$u_2 = -\frac{\alpha}{2\beta} + \frac{1}{2\beta}\sqrt{\frac{4\beta c - \alpha^2}{2}} \tanh\left(\sqrt{\frac{4\beta c - \alpha^2}{8\beta}}\xi\right),\tag{3.12}$$

where it requires that  $\alpha^2 - 4\beta c < 0$  and  $\beta > 0$ ; so these solutions must satisfy the condition

$$n^{2}g^{2}Q^{2} - 2\rho^{2}w^{2}N^{4} + \rho w N^{2}ngQ < 0, \qquad N^{2} > 0, \qquad (3.13)$$

which tells us that in the stable atmosphere  $(N^2 > 0)$ , heating (Q > 0) helps to generate downward-propagating waves, but cooling (Q < 0) helps to generate upward-propagating waves.



**Figure 1.** Propagation direction of solitary wave  $u_1$  (equation (3.9), where  $N^2 = 0.2, k = 1, n = -1, \omega = 2$ ) for (*a*) easterly Q = -0.1, (*b*) downward Q = -0.1, (*c*) westerly Q = 0.1, (*d*) upward Q = 0.1.

Equations (3.10) and (3.13) show that in the stable atmosphere ( $N^2 > 0$ ), heating (Q > 0) helps to generate easterly propagating waves if the slope of lines of constant phase is positive, and heating (Q > 0) helps to generate westerly propagating waves if the slope of lines of constant phase is negative.

However, cooling (Q < 0) helps to generate westerly propagating waves if the slope of lines of constant phase is positive, and cooling (Q < 0) helps to generate easterly propagating waves if the slope of lines of constant phase is negative:

$$u_3 = -\frac{\alpha}{2\beta} - \frac{\sqrt{6}}{\sqrt{-\beta}\xi}, \qquad \alpha^2 - 4\beta c = 0, \qquad \beta < 0,$$
 (3.14)

where it requires that  $n^2g^2Q^2 - 2\rho^2w^2N^4 + \rho wN^2ngQ = 0$  and  $N^2 < 0$ . In this case, there are no waves in unstable atmosphere:

$$u_{4} = -\frac{\alpha}{2\beta} + \frac{m}{2\beta} \sqrt{\frac{6(\alpha^{2} - 4\beta c)}{2m^{2} - 1}} cn\left(\sqrt{\frac{\alpha^{2} - 4\beta c}{4\beta(2m^{2} - 1)}}\xi\right),$$
(3.15)

where it requires that  $\alpha^2 - 4\beta c > 0$  and  $\beta > 0$ ; so these solutions must satisfy the condition

$$n^{2}g^{2}Q^{2} - 2\rho^{2}w^{2}N^{4} + \rho w N^{2}ngQ > 0, \qquad N^{2} > 0, \qquad (3.16)$$

which tells us that in the stable atmosphere  $(N^2 > 0)$ , heating (Q > 0) helps to generate upward-propagating periodic waves, but cooling (Q < 0) helps to generate downward-propagating periodic waves.

From equations (3.10) and (3.16) we see that in the stable atmosphere ( $N^2 > 0$ ), heating (Q > 0) helps to generate westerly propagating periodic waves if the slope of lines of constant phase is positive, and heating (Q > 0) helps to generate easterly propagating periodic waves if the slope of lines of constant phase is negative.

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**Figure 2.** Propagation of periodic wave  $u_4$  (equation (3.15), where  $N^2 = 0.2$ , k = 1, n = -1,  $\omega = 2$ ) for (a) easterly Q = -0.1, (b) downward Q = -0.1, (c) westerly Q = 0.1, (d) upward Q = 0.1.

However, cooling (Q < 0) helps to generate easterly propagating periodic waves if the slope of lines of constant phase is positive, and cooling (Q < 0) helps to generate westerly propagating periodic waves if the slope of lines of constant phase is negative. The corresponding propagating behaviour is given in figure 2:

$$u_{5} = -\frac{\alpha}{2\beta} + \frac{m}{2\beta} \sqrt{\frac{4\beta c - \alpha^{2}}{m^{2} + 1}} sn\left(\sqrt{\frac{4\beta c - \alpha^{2}}{4\beta(m^{2} + 1)}}\xi\right),$$
(3.17)

where it requires that  $\alpha^2 - 4\beta c < 0$  and  $\beta > 0$ ; so these solutions must satisfy the condition

$$n^{2}g^{2}Q^{2} - 2\rho^{2}w^{2}N^{4} + \rho w N^{2}ngQ < 0, \qquad N^{2} > 0, \qquad (3.18)$$

which tells us that in the stable atmosphere  $(N^2 > 0)$ , heating (Q > 0) helps to generate downward-moving periodic waves, but cooling (Q < 0) helps to generate upward-moving periodic waves.

From equations (3.10) and (3.18), it is not difficult to see that in the stable atmosphere  $(N^2 > 0)$ , heating (Q > 0) helps to generate easterly moving periodic waves if the slope of lines of constant phase is positive, and heating (Q > 0) helps to generate westerly moving periodic waves if the slope of lines of constant phase is negative. However, cooling (Q < 0) helps to generate westerly moving periodic waves if the slope of lines of constant phase is positive, and cooling (Q < 0) helps to generate easterly moving periodic waves if the slope of lines of constant phase is positive, and cooling (Q < 0) helps to generate easterly moving periodic waves if the slope of lines of constant phase is negative. In addition, when  $m \rightarrow 1$ ,  $u_4$  is reduced to  $u_1$  and  $u_5$  is reduced to  $u_2$ .

*Case 2.* If  $\gamma = 0$ , we set  $u = a_0 + \varphi$  and take equation (3.5) in the form

$$\frac{\mathrm{d}\varphi}{\mathrm{d}\xi} = \epsilon \sqrt{c_0 + c_2 \varphi^2 + c_4 \varphi^4},\tag{3.19}$$

which has the Weierstrass elliptic function solution

$$u_{6} = -\frac{\alpha}{2\beta} + \sqrt{\frac{\alpha^{2} - 4\beta c}{2\beta^{2}} - \frac{6}{\beta}} \wp(\xi; g_{2}, g_{3}), \qquad (3.20)$$

where

$$g_2 = \frac{4}{3} \left(\frac{\alpha^2}{4\beta} - c\right)^2 + \frac{2\beta c_0}{3}, \qquad g_3 = -\frac{8}{27} \left(\frac{3\alpha^2}{4\beta} - c\right)^3 - \frac{2\beta c_0}{9}, \quad (3.21)$$

$$u_7 = -\frac{\alpha}{2\beta} + 2\sqrt{3}\sqrt{\frac{c_0\beta}{12\wp\,(\xi;\,g_2,\,g_3)\beta - (\alpha^2 - 4\beta c)}},\tag{3.22}$$

in which  $g_2$  and  $g_3$  satisfy equation (3.21)

$$u_{8} = -\frac{\alpha}{2\beta} + \sqrt{\frac{12c_{0}\wp\left(\xi; g_{2}, g_{3}\right) + 2c_{0}\left(\frac{\alpha^{2} - 4c\beta}{2\beta} + D\right)}{12\wp\left(\xi; g_{2}, g_{3}\right) + D}},$$
(3.23)

where

$$g_2 = \frac{(-\alpha^2 + 4c\beta)(10D\beta + 2\alpha^2 - 8c\beta - 11c_0\beta^2)}{96\beta^2},$$
(3.24)

$$g_{3} = \frac{7D\alpha^{4}}{1152\beta^{2}} - \frac{7D\alpha^{2}c}{144\beta} + \frac{7Dc^{2}}{72} + \frac{7c_{0}\beta D}{144} + \frac{5\alpha^{6}}{3456\beta^{3}} - \frac{5\alpha^{4}c}{288\beta^{2}} + \frac{5\alpha^{2}c^{2}}{72\beta} - \frac{5c^{3}}{54} + \frac{c_{0}\alpha^{2}}{192} - \frac{cc_{0}\beta}{48}$$
(3.25)

$$D = \frac{5(-\alpha^2 + 4c\beta)}{8\beta} + \frac{1}{8}\sqrt{\frac{9(-\alpha^2 + 4c\beta)^2}{\beta^2} + 96\beta c_0}$$
(3.26)

$$u_{9} = -\frac{\alpha}{2\beta} - \frac{1}{12} \frac{\sqrt{c_{0}}(-24\wp(\xi; g_{2}, g_{3})\beta + 4c\beta - \alpha^{2})}{\beta\left(\frac{\partial}{\partial\xi}\wp(\xi; g_{2}, g_{3})\right)},$$
(3.27)

where

$$g_2 = \frac{(4c\beta - \alpha^2)^2}{192\beta^2} - \frac{c_0\beta}{6}, \qquad g_3 = \frac{(4c\beta - \alpha^2)}{864\beta} \left( 6c_0\beta + \frac{(4c\beta - \alpha^2)^2}{16\beta^2} \right), \tag{3.28}$$

$$u_{10} = -\frac{\alpha}{2\beta} + \frac{12\sqrt{6}\sqrt{-\beta}\frac{\partial}{\partial\xi}\wp(\xi;g_2,g_3)}{24\wp(\xi;g_2,g_3)\beta - 4c\beta + \alpha^2},$$
(3.29)

in which  $g_2$  and  $g_3$  satisfy equation (3.28). In this case,  $\wp(\xi; g_2, g_3)$  in the solutions  $u_6$  to  $u_{10}$  is the Weierstrass elliptic function.

*Case 3.* If  $\gamma \neq 0$ , when we make a transformation  $\varphi \rightarrow \varphi^{2/3}$  for equation (3.5), then equation (3.5) is reduced to

$$\frac{\mathrm{d}\varphi}{\mathrm{d}\xi} = \frac{3}{2}\epsilon\sqrt{c_0\varphi^{2/3} + c_1\varphi^{4/3} + c_2\varphi^2 + c_3\varphi^{8/3} + c_4\varphi^{10/3} + c_5\varphi^4}.$$
(3.30)

Suppose that  $c_0 = c_1 = c_3 = c_4 = 0$ , and thus equation (3.1) possesses a bell-shaped soliton solution

$$u_{11} = -\frac{\beta}{3\gamma} + \left(\frac{10(2\beta^3 + 27c\gamma^2 - 9\alpha\beta\gamma)}{27\gamma^3}\operatorname{sech}^2\left(\frac{\sqrt{-6\beta^3 - 81c\gamma^2 + 27\alpha\beta\gamma}}{6\gamma}\xi\right)\right)^{1/3}, \quad (3.31)$$



**Figure 3.** The states of a triangular wave solution  $u_{12}$  (equation (3.33), where  $N^2 = 0.2, \omega = 2, n = 1, t = 0$ ) at the *x*-*z* plane for (*a*) Q = 0.1, k = -1, (b) Q = -0.1, k = 1.

where it requires that  $2\beta^3 - 9\gamma\alpha\beta + 27c\gamma^2 < 0$ ; so this solution must satisfy the condition

$$5\rho\omega^2 N^2 - 4gn\omega Q < 0, \tag{3.32}$$

which tells us that heating (Q > 0) helps to generate upward-propagating solitary waves, but cooling (Q < 0) helps to generate downward-propagating solitary waves.

From equations (3.10) and (3.32), we see that heating (Q > 0) helps to generate westerly propagating solitary waves if the slope of lines of constant phase is positive, and heating (Q > 0) helps to generate easterly propagating solitary waves if the slope of lines of constant phase is negative. However, cooling (Q < 0) helps to generate easterly propagating solitary waves if the slope of lines of constant phase is positive, and cooling (Q < 0) helps to generate westerly propagating solitary waves if the slope of lines of constant phase is positive, and cooling (Q < 0) helps to generate westerly propagating solitary waves if the slope of lines of constant phase is negative.

A triangular solution

$$u_{12} = -\frac{\beta}{3\gamma} + \left(\frac{10(2\beta^3 + 27c\gamma^2 - 9\alpha\beta\gamma)}{27\gamma^3}\sec^2\left(\frac{\sqrt{6\beta^3 + 81c\gamma^2 - 27\alpha\beta\gamma}}{6\gamma}\xi\right)\right)^{1/3}, \quad (3.33)$$

where it requires that  $2\beta^3 - 9\gamma\alpha\beta + 27c\gamma^2 > 0$ ; so this solution must satisfy the condition

$$5\rho\omega^2 N^2 - 4gn\omega Q > 0, \tag{3.34}$$

which tells us that heating (Q > 0) helps to generate downward-propagating waves, but cooling (Q < 0) helps to generate upward-propagating waves.

From equations (3.10) and (3.34), it is not difficult to see that heating (Q > 0) helps to generate easterly propagating waves if the slope of lines of constant phase is positive, and heating (Q > 0) helps to generate westerly propagating waves if the slope of lines of constant phase is negative. However, cooling (Q < 0) helps to generate westerly propagating waves if the slope of lines of constant phase is positive, and cooling (Q < 0) helps to generate easterly propagating waves if the slope of lines of constant phase is positive, and cooling (Q < 0) helps to generate easterly propagating waves if the slope of lines of constant phase is negative. The corresponding propagating behaviour is given in figures 3 and 4.

A rational solution

$$u_{13} = -\frac{\beta}{3\gamma} \pm \left(\frac{10}{\gamma\xi^2}\right)^{1/3},$$
(3.35)

where it requires that  $5\rho\omega^2 N^2 - 4gn\omega Q = 0$ , which shows that the heat term have nearly no effect on the generation and propagation of internal gravity waves.



**Figure 4.** Propagation of a triangular wave solution  $u_{12}$  (equation (3.33), where Q = -0.1,  $N^2 = 0.2$ ,  $\omega = 2$ , n = 1) for (a) westerly (k = -1), (b) upward (k = 1), (c) easterly (k = 1), (d) upward (k = -1).

When we take s = 3 and r = 5 in equation (1.3), namely making a transformation  $\varphi \rightarrow \varphi^{1/3}$  for equation (3.5), then equation (3.5) is reduced to

$$\frac{\mathrm{d}\varphi}{\mathrm{d}\xi} = 3\epsilon\sqrt{c_0\varphi^{4/3} + c_1\varphi^{5/3} + c_2\varphi^2 + c_3\varphi^{7/3} + c_4\varphi^{8/3} + c_5\varphi^3}.$$
(3.36)

Suppose that  $c_0 = c_1 = c_3 = c_4 = 0$ ; then equation (3.1) also have the same solitary wave solutions as  $u_{11}$ ,  $u_{12}$  and  $u_{13}$ .

#### 3.2. The first-order ODE with an eighth-degree nonlinear term

With the aid of Maple9, substitution of equation (3.6) into equation (3.1) shows that the set of algebraic equations possesses the solutions, as detailed in the following sections.

*3.2.1. The first kind of solutions with a fifth-degree nonlinear term.* The first set of solutions reads

 $c_{0} = c_{0}, \qquad c_{1} = c_{1}, \qquad c_{2} = -c - \beta a_{0}^{2} - \gamma a_{0}^{3} - \alpha a_{0},$   $c_{3} = -\gamma a_{0}^{2} a_{1} - \frac{2\beta a_{0} a_{1}}{3} - \frac{\alpha a_{1}}{3}, \qquad c_{4} = -\frac{\gamma a_{0} a_{1}^{2}}{2} - \frac{\beta a_{1}^{2}}{6}, \qquad c_{5} = -\frac{\gamma a_{1}^{3}}{10},$   $c_{6} = 0, \qquad c_{7} = 0, \qquad c_{8} = 0, \qquad a_{0} = a_{0}, \qquad a_{1} = a_{1}, \qquad a_{2} = 0,$ (3.37)

where  $a_1 \neq 0$ ,  $a_0$ , c and  $c_0$  are arbitrary constants. In this case, its solutions have been studied in section 3.1.

*3.2.2. The second kind of solutions with an eighth-degree nonlinear term.* The second set of solutions reads

$$c_{0} = 0, c_{1} = 0, c_{2} = -\frac{c}{4} - \frac{\beta a_{0}^{2}}{4} - \frac{\gamma a_{0}^{3}}{4} - \frac{\alpha a_{0}}{4}, c_{3} = 0,$$

$$c_{4} = -\frac{\beta a_{0} a_{2}}{6} - \frac{\gamma a_{0}^{2} a_{2}}{4} - \frac{\alpha a_{2}}{12}, c_{5} = 0, c_{6} = -\frac{\beta a_{2}^{2}}{24} - \frac{\gamma a_{0} a_{2}^{2}}{8}, c_{7} = 0,$$

$$c_{8} = -\frac{\gamma a_{2}^{3}}{40}, a_{0} = a_{0}, a_{1} = 0, a_{2} = a_{2}, (3.38)$$

where  $a_i (i = 0, 1, 2)$  and  $c_j (j = 0, 1, ..., 8)$  are arbitrary constants.

Then equation (3.6) takes in the form

$$\frac{\mathrm{d}\varphi}{\mathrm{d}\xi} = \epsilon \sqrt{c_2 \varphi^2 + c_4 \varphi^4 + c_6 \varphi^6 + c_8 \varphi^8}.$$
(3.39)

Let us make a transformation  $\varphi \rightarrow \varphi^{1/2}$  for equation (3.39); then equation (3.39) is reduced to

$$\frac{d\varphi}{d\xi} = 2\epsilon \sqrt{c_2 \varphi^2 + c_4 \varphi^3 + c_6 \varphi^4 + c_8 \varphi^5}.$$
(3.40)

If  $\gamma = 0$  and we choose  $a_0 = 0$ , then  $c_2 = -c$ ,  $c_4 = -\frac{\alpha a_1}{3}$ ,  $c_6 = -\frac{\beta a_1^2}{6}$ , and taking equation (3.40) in the form

$$\frac{\mathrm{d}\varphi}{\mathrm{d}\xi} = \epsilon \sqrt{c_2 \varphi^2 + c_4 \varphi^3 + c_6 \varphi^4} \tag{3.41}$$

which yields the following solitary wave solutions and triangular periodic wave solutions

$$u_{14} = \frac{6c\alpha\operatorname{sech}^2(\frac{\sqrt{-c}\xi}{2}\xi)}{-2\alpha^2 + 3c\beta - 6c\beta\operatorname{tanh}(\frac{\sqrt{-c\xi}}{2}) + 3c\beta\operatorname{tanh}^2(\frac{\sqrt{-c\xi}}{2})}, \qquad c < 0,$$
(3.42)

where it must satisfy  $Qng\omega - \rho\omega^2 N^2 > 0$ , it shows that heating helps to generate upwardpropagating waves, but cooling helps to generate downward-propagating waves. If the slope of lines of constant phase is positive, heating helps to generate westerly propagating waves, and if the slope of lines of constant phase is negative, heating helps to generate easterly propagating waves. However, cooling helps to generate easterly propagating waves if the slope of lines of constant phase is positive, and cooling helps to generate westerly propagating waves if the slope of lines of constant phase is negative:

$$u_{15} = \mp \frac{6c \operatorname{sech}(\sqrt{-c\xi})}{\sqrt{(\alpha^2 - 6c\beta)} \pm \alpha \operatorname{sech}(\sqrt{-c\xi})}, \qquad \alpha^2 - 6c\beta > 0, \qquad c < 0$$
(3.43)

$$u_{16} = -\frac{6c \operatorname{sech}^2\left(\frac{\sqrt{-c\xi}}{2}\right)}{\pm 2\sqrt{\alpha^2 - 6c\beta} \mp (\sqrt{\alpha^2 - 6c\beta} \pm \alpha) \operatorname{sech}^2\left(\frac{\sqrt{-c\xi}}{2}\right)}, \qquad \alpha^2 - 6c\beta > 0, \qquad c < 0,$$
(3.44)

$$u_{17} = -\frac{6c\operatorname{csch}^2\left(\frac{\sqrt{-c\xi}}{2}\right)}{\pm 2\sqrt{\alpha^2 - 6c\beta} \pm (\sqrt{\alpha^2 - 6c\beta} \pm \alpha)\operatorname{csch}^2\left(\frac{\sqrt{-c\xi}}{2}\right)}, \qquad \alpha^2 - 6c\beta > 0, \qquad c < 0,$$

, \_

where the solutions  $(u_{15}, u_{16} \text{ and } u_{17})$  must satisfy the condition  $2Q^2g^2n^2 + 5QgnN^2\rho\omega - 7N^4\rho^2\omega^2 > 0, \qquad Qng\omega - \rho\omega^2N^2 > 0,$ (3.46)



**Figure 5.** Propagation direction of a solitary wave  $u_{17}$  (equation (3.45), where Q = 0.1, n = 1,  $N^2 = 0.2, \omega = 2$ ) for (*a*) westerly (k = -1), (*b*) upward (k = 1), (*c*) easterly (k = 1), (*d*) upward (k = -1).

which tells us that heating tends to generate upward-moving waves, but cooling tends to generate downward-moving waves. If the slope of lines of constant phase is positive, heating tends to generate westerly moving waves, and if the slope of lines of constant phase is negative, heating tends to generate easterly moving waves. However, cooling tends to generate easterly moving waves. However, cooling tends to generate easterly moving waves if the slope of lines of constant phase is positive, and cooling helps to generate westerly moving waves if the slope of lines of constant phase is negative. Figure 5 shows the propagating behaviour of the positive branch of  $u_{17}$  in qualitative terms

$$u_{18} = -\frac{6c\sec^2\left(\frac{\sqrt{c\xi}}{2}\right)}{\pm 2\sqrt{\alpha^2 - 6c\beta} \mp (\sqrt{\alpha^2 - 6c\beta} \pm \alpha)\sec^2\left(\frac{\sqrt{c\xi}}{2}\right)}, \qquad \alpha^2 - 6c\beta > 0, \qquad c > 0,$$
(3.47)

$$u_{19} = -\frac{6c\csc^2\left(\frac{\sqrt{c\xi}}{2}\right)}{\pm 2\sqrt{\alpha^2 - 6c\beta} \pm (\sqrt{\alpha^2 - 6c\beta} \pm \alpha)\csc^2\left(\frac{\sqrt{c\xi}}{2}\right)}, \qquad \alpha^2 - 6c\beta > 0, \qquad c > 0,$$
(3.48)

where the solutions  $(u_{18} \text{ and } u_{19})$  must satisfy the condition

$$2Q^{2}g^{2}n^{2} + 5QgnN^{2}\rho\omega - 7N^{4}\rho^{2}\omega^{2} > 0, \qquad Qng\omega - \rho\omega^{2}N^{2} < 0, \qquad (3.49)$$

which tells us that under the influence of heating, wave fronts tend to be travelling downward; if the slope of lines of constant phase is positive, wave fronts tend to be travelling in a positive *x*-direction; and if the slope of lines of constant phase is negative, wave fronts tend to be travelling in a negative *x*-direction. However, under the influence of cooling, wave fronts tend to be travelling upward; if the slope of lines of constant phase is positive, wave fronts tend to be travelling in a negative *x*-direction; and if the slope of lines of constant phase is positive, wave fronts tend to be travelling in a negative *x*-direction; and if the slope of lines of constant phase is positive, wave fronts tend to be travelling in a negative *x*-direction; and if the slope of lines of constant phase is negative,



**Figure 6.** Propagation of wave  $u_{19}$  (equation (3.48), where Q = -0.1, n = 1,  $N^2 = 0.2$ ,  $\omega = 2$ ) for (*a*) westerly (k = -1), (*b*) upward (k = 1), (*c*) easterly (k = 1), (*d*) upward (k = -1).

wave fronts tend to be travelling in a positive x-direction. Figure 6 shows the propagating behaviour of the positive branch of  $u_{19}$  in qualitative terms.

If  $\gamma \neq 0$ , when we make a transformation  $\varphi \rightarrow \varphi^{2/3}$  for equation (3.40), then equation (3.40) is reduced to

$$\frac{\mathrm{d}\varphi}{\mathrm{d}\xi} = \frac{3}{2}\epsilon\sqrt{c_2\varphi^2 + c_4\varphi^{8/3} + c_6\varphi^{10/3} + c_8\varphi^4}.$$
(3.50)

Suppose that  $c_4 = c_6 = 0$ ; then equation (3.1) also have the same solitary wave solutions as  $u_{11}$ ,  $u_{12}$  and  $u_{13}$ .

3.2.3. The third kind of solutions with an eighth-degree nonlinear term. The third set of solutions reads

$$c_{1} = -a_{1} \left(-320\beta a_{0}a_{1}^{2}a_{2}^{2} + 20\beta a_{1}^{4}a_{2} + 60\gamma a_{0}a_{1}^{4}a_{2} - 3\gamma a_{1}^{6} - 480\gamma a_{0}^{2}a_{1}^{2}a_{2}^{2} - 160a_{1}^{2}a_{2}^{2}\alpha + 1920ca_{2}^{3} + 1920\gamma a_{0}^{3}a_{2}^{3} + 1920\beta a_{0}^{2}a_{2}^{3} + 1920\alpha a_{0}a_{2}^{3}\right) / (7680a_{2}^{4}),$$

$$c_{2} = -\left(320\beta a_{0}a_{1}^{2}a_{2}^{2} - 20\beta a_{1}^{4}a_{2} - 60\gamma a_{0}a_{1}^{4}a_{2} + 3\gamma a_{1}^{6} + 480\gamma a_{0}^{2}a_{1}^{2}a_{2}^{2} + 160a_{1}^{2}a_{2}^{2}\alpha + 640ca_{2}^{3} + 640\gamma a_{0}^{3}a_{2}^{3} + 640\beta a_{0}^{2}a_{2}^{3} + 640\alpha a_{0}a_{2}^{3}\right) / (2560a_{2}^{3}),$$

$$c_{3} = -\frac{a_{1}\left(320\beta a_{0}a_{2}^{2} + 20\beta a_{1}^{2}a_{2} + 60\gamma a_{0}a_{1}^{2}a_{2} - 3\gamma a_{1}^{4} + 480\gamma a_{0}^{2}a_{2}^{2} + 160\alpha a_{2}^{2}\right)}{960a_{2}^{2}},$$

$$c_{4} = -\frac{64\beta a_{0}a_{2}^{2} + 44\beta a_{1}^{2}a_{2} + 132\gamma a_{0}a_{1}^{2}a_{2} + 3\gamma a_{1}^{4} + 96\gamma a_{0}^{2}a_{2}^{2} + 32\alpha a_{2}^{2}}{384a_{2}},$$

$$c_{5} = -\frac{\beta a_{1}a_{2}}{8} - \frac{13\gamma a_{1}^{3}}{160} - \frac{3\gamma a_{0}a_{1}a_{2}}{8}, \qquad c_{6} = -\frac{\beta a_{2}^{2}}{24} - \frac{23\gamma a_{1}^{2}a_{2}}{160} - \frac{\gamma a_{0}a_{2}^{2}}{8},$$

$$13$$

$$c_{7} = -\frac{\gamma a_{1} a_{2}^{2}}{10}, \qquad c_{8} = -\frac{\gamma a_{2}^{3}}{40}, \qquad c_{0} = c_{0},$$
  
$$a_{0} = a_{0}, \qquad a_{1} = a_{1}, \qquad a_{2} = a_{2}, \qquad (3.51)$$

where  $a_i$  (i = 0, 1, 2) and  $c_j$  (j = 0, 1, ..., 8) are arbitrary constants. Considering the 11 parameters, we present some types of solutions in the following cases.

*Case 1.* If  $c_7 = c_8 = 0$  in equation (3.6), then equation (3.6) becomes

$$\frac{\mathrm{d}\varphi}{\mathrm{d}\xi} = \epsilon \sqrt{c_0 + c_1 \varphi + c_2 \varphi^2 + c_3 \varphi^3 + c_4 \varphi^4 + c_5 \varphi^5 + c_6 \varphi^6}.$$
(3.52)

By using the condition  $c_7 = c_8 = 0$ , we can obtain that  $a_1 = 0$ ,  $\gamma = 0$ ,  $c_1 = c_3 = c_5 = 0$ . Suppose that  $c_0 = \frac{8c_2^2}{27c_4}$  and  $c_6 = \frac{c_4^2}{4c_2}$ ; then we can obtain  $a_0 = \frac{-\alpha \pm \sqrt{3\alpha^2 - 12c\beta}}{2\beta}$ . Then equation (3.1) has a kink profile solution

$$u_{20} = \frac{-\alpha \pm \sqrt{3\alpha^2 - 12c\beta}}{2\beta} \pm \frac{4(\alpha^2 - 4\beta c) \tanh^2(\pm \frac{1}{12}\sqrt{\frac{6\alpha^2 - 24\beta c}{\beta}}\xi)}{\beta\sqrt{3\alpha^2 - 12\beta c}(3 + \tanh^2(\pm \frac{1}{12}\sqrt{\frac{6\alpha^2 - 24\beta c}{\beta}}\xi))}$$
(3.53)

and a singular solution

$$u_{21} = \frac{-\alpha \pm \sqrt{3\alpha^2 - 12c\beta}}{2\beta} \pm \frac{4(\alpha^2 - 4\beta c) \coth^2(\pm \frac{1}{12}\sqrt{\frac{6\alpha^2 - 24\beta c}{\beta}}\xi)}{\beta\sqrt{3\alpha^2 - 12\beta c}(3 + \coth^2(\pm \frac{1}{12}\sqrt{\frac{6\alpha^2 - 24\beta c}{\beta}}\xi))},$$
(3.54)

where  $\alpha^2 - 4\beta c > 0$  and  $\beta > 0$  in equations (3.53) and (3.54); so the solutions must satisfy condition (3.10), where the wave propagation directions have been studied in detail in section (3.1).

A triangular periodic solution

$$u_{22} = \frac{-\alpha \pm \sqrt{3\alpha^2 - 12c\beta}}{2\beta} \pm \frac{4(\alpha^2 - 4\beta c)\tan^2\left(\pm\frac{1}{12}\sqrt{-\frac{6\alpha^2 - 24\beta c}{\beta}}\xi\right)}{\beta\sqrt{3\alpha^2 - 12\beta c}\left(3 - \tan^2\left(\pm\frac{1}{12}\sqrt{-\frac{6\alpha^2 - 24\beta c}{\beta}}\xi\right)\right)}$$
(3.55)

and a singular triangular periodic solution

$$u_{23} = \frac{-\alpha \pm \sqrt{3\alpha^2 - 12c\beta}}{2\beta} \pm \frac{4(\alpha^2 - 4\beta c)\cot^2(\pm \frac{1}{12}\sqrt{-\frac{6\alpha^2 - 24\beta c}{\beta}}\xi)}{\beta\sqrt{3\alpha^2 - 12\beta c}(3 - \cot^2(\pm \frac{1}{12}\sqrt{-\frac{6\alpha^2 - 24\beta c}{\beta}}\xi))},$$
(3.56)

where  $\alpha^2 - 4\beta c > 0$  and  $\beta < 0$  in equations (3.55) and (3.56) respectively; so the solutions must satisfy the condition

$$Q^{2}g^{2}n^{2} + QgnN^{2}\rho\omega - 2N^{4}\rho^{2}\omega^{2} > 0, \qquad N^{2} < 0, \qquad (3.57)$$

which tells us that in the unstable atmosphere  $(N^2 < 0)$ , heating (Q > 0) helps to generate downward-propagating waves, but cooling (Q < 0) helps to generate upward-propagating waves.

From equations (3.10) and (3.57), it is not difficult to see that in the unstable atmosphere  $(N^2 < 0)$ , heating (Q > 0) helps to generate easterly propagating waves if the slope of lines of constant phase is positive, and heating (Q > 0) helps to generate westerly propagating waves if the slope of lines of constant phase is negative.

However, in the unstable atmosphere  $(N^2 < 0)$ , cooling (Q < 0) helps to generate westerly propagating waves if the slope of lines of constant phase is positive, and cooling

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Table 1. The wave propagation direction from the local heating.						
Solution type	Stability	Upward	Downward	Westerly	Easterly	ODE type
$u_1, u_4$	$N^{2} > 0$	Q > 0	Q < 0	$Q \cdot Sp > 0$	$Q \cdot Sp < 0$	ODE5, ODE8
$u_2, u_5$	$N^{2} > 0$	Q < 0	Q > 0	$Q \cdot Sp < 0$	$Q \cdot Sp > 0$	ODE5, ODE8
<i>u</i> <sub>11</sub>	_	Q > 0	Q < 0	$Q \cdot Sp > 0$	$Q \cdot Sp < 0$	ODE5, ODE8
<i>u</i> <sub>12</sub>	_	Q < 0	Q > 0	$Q \cdot Sp < 0$	$Q \cdot Sp > 0$	ODE5, ODE8
$u_{14}, u_{15}, u_{16}, u_{17}$	_	Q > 0	Q < 0	$Q \cdot Sp > 0$	$Q \cdot Sp < 0$	ODE8
$u_{18}, u_{19}$	_	Q < 0	Q > 0 -	$Q \cdot Sp < 0$	$Q \cdot Sp > 0$	ODE8
$u_{22}, u_{23}$	$N^2 < 0$	Q < 0	Q > 0	$Q \cdot Sp < 0$	$Q \cdot Sp > 0$	ODE8
1120 1121 1124 1125	$N^2 > 0$	Q > 0	Q < 0	$Q \cdot Sp > 0$	$Q \cdot Sp < 0$	ODE8

(Q < 0) helps to generate easterly propagating waves if the slope of lines of constant phase is negative.

If we suppose that  $h_0 = 0$  and  $c_6 = \frac{c_4^2}{4c_2}$ , then equation (3.1) has a kink profile solution

$$u_{24} = \frac{-\alpha \pm \sqrt{3\alpha^2 - 12c\beta}}{2\beta} \pm \frac{3(\alpha^2 - 4\beta c)}{2\beta\sqrt{3\alpha^2 - 12\beta c}} \left(1 + 3\tanh\left(\pm\sqrt{\frac{2\alpha^2 - 8\beta c}{\beta}}\xi\right)\right)$$
(3.58)

and a singular solution

$$u_{25} = \frac{-\alpha \pm \sqrt{3\alpha^2 - 12c\beta}}{2\beta} \pm \frac{3(\alpha^2 - 4\beta c)}{2\beta\sqrt{3\alpha^2 - 12\beta c}} \left(1 + 3\coth\left(\pm\sqrt{\frac{2\alpha^2 - 8\beta c}{\beta}}\xi\right)\right), \quad (3.59)$$

where  $\alpha^2 - 4\beta c > 0$  and  $\beta > 0$  in equations (3.58) and (3.59) respectively; so the solutions must satisfy condition (3.10), where the wave propagation directions have been studied in detail in section (3.1).

*Case 2.* If we take s = 5 and r = 8 in equation (1.3), namely making a transformation  $\varphi \rightarrow \varphi^{1/2}$  for equation (3.6), then equation (3.6) becomes

$$\frac{\mathrm{d}\varphi}{\mathrm{d}\xi} = 2\epsilon\sqrt{c_0\varphi + c_1\varphi^{3/2} + c_2\varphi^2 + c_3\varphi^{5/2} + c_4\varphi^3 + c_5\varphi^{7/2} + c_6\varphi^4 + c_7\varphi^{9/2} + c_8\varphi^5}.$$
 (3.60)

Supposing that  $c_1 = c_3 = c_5 = c_7 = 0$ , equation (3.1) also has the same solitary wave solutions as  $u_{11}$ ,  $u_{12}$  and  $u_{13}$ .

#### 4. Summary and discussions

We have extended Fan's direct and unified algebraic method with symbolic computation to the case when r > 4 by introducing a new transformation and successfully applied this to investigate internal gravity waves in the atmosphere. The solutions of a first-order nonlinear ordinary differential equation with a higher degree nonlinear term, such as with a fifth-degree nonlinear term (ODE5) and an eighth degree nonlinear term (ODE8), are obtained, which included periodic wave solutions associated with the Jacobin elliptic function and the bellshaped and kink profile solitary wave solutions. Most importantly, the propagation and generation of gravity waves are affected by the local heating conditions; the results are listed in table 1, where  $Sp = \frac{dz}{dx}$  denotes the slope of lines of constant phase in the *x*-*z* plane,  $Q \cdot Sp > 0$  denotes Q > 0 and Sp > 0 or Q < 0 and Sp < 0, and  $Q \cdot Sp < 0$  denotes Q > 0and Sp < 0 or Q < 0 and Sp > 0. Except those considered in this paper, the proposed method is of great significance in many fields in physics, mechanics, atmosphere and ocean, etc, and it is also readily applicable to a large variety of other nonlinear evolution equations such as the generalized coupled Hirota–Satsuma, coupled Schrödinger–KdV, (2+1)-dimensional dispersive long wave, (2+1)-dimensional Davey–Stewartson equations, the (3+1)-dimensional Jimbo–Miwa equation, etc. The details for these cases will be investigated in our future work.

# Acknowledgments

We would like to express our sincere thanks to the referees for their useful comments. This work was supported by the National Natural Science Foundation of China (grant no. 40775069).

### References

- [1] Smith R B and Lin Y L 1982 Q. J. R. Meteorol. Soc. 108 353
- [2] Li Z L 2007 Chin. J. Geophys. 50 34
- [3] Lin Y L and Chun H Y 1991 J. Atmos. Sci. 48 2476
- [4] Fritts D C and Nastrom G D 1992 J. Atmos. Sci. 49 111
- [5] Drazin P G and Johnson R S 1989 Solitons: An Introduction (Cambridge: Cambridge University Press)
- [6] Wadati M 1975 J. Phys. Soc. Japan. 38 673
- [7] Wadati M 1975 J. Phys. Soc. Japan. 38 681
- [8] Konno K 1974 J. Phys. Soc. Japan. 37 1631
- [9] Otwinowski M, Paul R and Laidlaw W G 1998 Phys. Lett. A 128 483
- [10] Mohamad M N 1992 Math. Methods Appl. Sci. 15 73
- [11] Coffey M W 1990 SIAM J. Appl. Math. 50 1580
- [12] Ablowitz M J and Clarkson P A 1991 Soliton, Nonlinear Evolution Equations and Inverse Scattering (Cambridge: Cambridge University Press)
- [13] Miurs M R 1978 Bäcklund Transformation (Berlin: Springer)
- [14] Lamb G L 1971 Rev. Mod. Phys. 43 99
- [15] Hirota R 1973 J. Math. Phys. 14 810
- [16] Hirota R and Satsuma J 1981 Phys. Lett. A 85 407
- [17] Satsuma J and Hirota R 1982 J. Phys. Soc. Japan 51 332
- [18] El-Sabbagh M F and Ali A T 2005 Int. J. Nonlinear Sci. Numer. Simul. 6 151
- [19] Abdusalam H A 2005 Int. J. Nonlinear Sci. Numer. Simul. 6 99
- [20] Parkes E J and Duffy B R 1996 Comput. Phys. Commun. 98 288
- [21] Parkes E J and Duffy B R 1997 Phys. Lett. A 229 217
- [22] Duffy B R and Parkes E J 1996 Phys. Lett. A 214 271
- [23] Hereman W 1991 Comput. Phys. Commun. 65 143
- [24] Hereman W and Nuseir A 1997 Math. Comput. Simul. 43 13
- [25] Fan E G 2000 Phys. Lett. A 35 6853
- [26] Fan E G 2001 Z. Naturforsch. A 56 312
- [27] Abdou M A and Soliman A A 2005 Physica D 211 1
- [28] Abdou M A and Soliman A A 2005 J. Comput. Appl. Math. 181 245
- [29] Soliman A A and Raslan K R 2001 Int. J. Comput. Math. 78 399
- [30] Soliman A A 2004 Int. J. Comput. Math. 81 325
- [31] Soliman A A and Hussein M H 2005 Appl. Math. Comput. 161 623
- [32] Abassy T A, El-Tawil M A and Saleh H K 2004 Int. J. Nonlinear Sci. Numer. Simul. 5 327
- [33] Beals R and Coifman R R 1984 Commun. Pure. Appl. Math. 37 39
- [34] Matveev V B and Salle M A 1991 Darboux Transformation and Solitons (Berlin: Springer)
- [35] Leble S B and Ustinov N V 1993 J. Phys. A: Math. Gen. 26 5007
- [36] Esteevez P G 1999 J. Math. Phys. 40 1406
- [37] Dubrousky V G and Konopelchenko B G 1994 J. Phys. A: Math. Gen. 27 4619
- [38] Neugebauer G and Kramerl D 1983 J. Phys. A: Math. Gen. 16 1927
- [39] Wang M L, Zhou Y B and Li Z B 1996 Phys. Lett. A 216 67
- [40] Yang L, Zhou Z and Wang Y 1999 Phys. Lett. A 260 55

- [41] Sirendaoreji J S 2003 Phys. Lett. A **309** 387
- [42] Yomba E 2004 Chaos Solitons and Fractals 22 321
- [43] Yomba E 2004 Chaos Solitons and Fractals 21 75
- [44] Fan E G 2003 *Phys. Lett.* A **300** 243
- [45] Fan E G 2002 J. Phys. A: Math. Gen. 35 6853
- [46] Sirendaoreji J S 2006 Phys. Lett. A 356 124
- [47] Belokolos E, Bobenko A, Enol'skij V, Its A and Matveev V 1994 Algebro-Geometrical Approach to Nonlinear Integral Equations (Berlin: Springer)
- [48] Christiansen P L, Eilbrck J C, Enolskii V Z and Kostov N A 1995 Proc. R. Soc. A 451 685
- [49] Samsonov A M 1995 Appl. Anal. 57 85
- [50] Porubov A V and Paeker D F 1999 Wave Motion 29 97
- [51] Akhiezer N L 1990 Elements of Theory of Elliptic Functions (Providence, RI: American Mathematical Society)
- [52] Matsuno Y 2001 J. Math. Phys. 42 1744
- [53] Maccari A 1998 J. Math. Phys. 35 1755
- [54] Singh S V, Rao N N and Shukla P K 1998 J. Plasma Phys. 60 551
- [55] Turner J S 1973 Buoyancy Effects in Fluids (Cambridge: Cambridge University Press)
- [56] Nappo C J 2002 An Introduction to Atmospheric Gravity Waves (London: Academic)